

A fracture-based discrete model for simulating creep in quartz sands

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Abstract (150 – 200 words)

Creep of granular soils is frequently accompanied by grain breakage. Stress corrosion driven grain breakage offers a micromechanically based explanation for granular creep. This study incorporates that concept into a new model based on the discrete element method (DEM) to simulate creep in sands. The model aims for conceptual simplicity, computational efficiency and ease of calibration. To this end a new form of normalized Charles power law is incorporated into a DEM model for rough-crushable sands based on the particle splitting technique. The model is implemented using a controlled on-off computational strategy. The model is validated by simulating creep in quartz sands in oedometric and triaxial conditions. Model predictions are shown to compare favourably with experimental results in terms of creep strain, creep strain rates and particle breakage. The model proposed would facilitate the calibration of phenomenological continuum models, but may be also useful to directly investigate structural scale phenomena, such as pile ageing.

Keywords chosen from ICE Publishing list

Cracks & cracking; Creep; Discrete-element modelling; Particle crushing; Time dependence;

List of notations (examples below)

- δ is the particle overlap
- δ_{T1}, δ_{T2} are critical contact overlaps for three transition regimes
- δ_1, δ_2 are the dimensional constants
- F_n is the contact normal force
- F_{nT1}, F_{nT2} is the transitional contact normal forces
- E_1, E_2 are young's modulus of two contacting particles
- ν_1, ν_2 are Poisson's ratios of two contacting particles
- r_1, r_2 are radii of two contacting particles
- S_q is the particle roughness
- A_F is the force contact area
- κ_{mob}, κ are the mobilised and intrinsic strengths
- χ is the microstructure parameter
- ν is the Poisson's ratio
- R is the particle radius
- σ_{lim} is the characteristic particle limiting stress
- $\sigma_{lim,0}$ is the mean particle limiting stress
- d is the particle diameter
- d_0 is the reference particle diameter
- m_p is the material parameter
- M_T is the total mass

d_{max}, d_{min} are the maximum//minimum particle diameter
 β is the fractal factor
 v is the crack velocity
 v_0 is the reference velocity
 K_c is the material toughness
 n is the stress corrosion index
 β_c is the crack geometry parameter
 a is the crack half length
 σ_t is the tensile strength
 σ_{mob} is the particle maximum normal contact stress
 m is the creep strain rate parameter
 a_0 is the initial crack half-length

1 Introduction

2 The time-dependent behaviour of granular soils has been recognized as a crucial mechanism in
3 many phenomena of engineering significance like the observed large shaft capacity increase with
4 time of driven piles, known as pile set-up (Jardine *et al.*, 2006; Zhang & Wang, 2015; Gavin &
5 Igoe, 2021). From the material viewpoint, pile set-up is a particular manifestation of ageing
6 (Schmertmann, 1991), a process by which the mechanical properties of soils improve with time
7 without significant change in effective stress. It is currently recognized (Mitchell, 2008) that the
8 most plausible explanation of sand ageing is given by physically-driven -as opposed to chemically
9 or biologically mediated- micromechanical evolution. Physically-driven micromechanical evolution
10 would necessarily involve changes in the granular contact fabric around the foundation (Bowman
11 & Soga, 2003). However, absent any external disturbance, time changes in granular fabric can
12 only take place if there are physical changes in the grains themselves, i.e. in their sizes, shapes
13 or contact properties.

14
15 Laboratory studies have identified some of the physical modifications that underlie time effects in
16 sand. It is now clear that grain crushing is associated with time dependent phenomena in sands
17 (Colliat-Dangus *et al.*, 1988; Takei *et al.*, 2001; Karimpour & Lade, 2010; Brzesowsky *et al.*, 2014;
18 Lv *et al.*, 2017). Direct evidence of grain fracture during experiments may be obtained with optical
19 imaging techniques (e.g. Takei *et al.*, 2001) or by means of computed tomography (Andó *et al.*,
20 2019). More frequently, evidence for grain crushing is only obtained after the test is finished, by
21 examining grain size distributions (GSD). This experimental procedure is not only slow and
22 inefficient (one measurement point per sample) but also prone to error when materials are too
23 friable or the fragments are too small (Karimpour & Lade, 2010). Physical modifications related to
24 time effects are also present at the sub-granular scale. Asperity and roughness evolve under
25 loading, leading to what is described as contact maturing (Michalowski *et al.*, 2018). It is likely
26 that contact maturing is more relevant at lower stress levels than grain fracture. Needless to say,
27 the experimental difficulties only increase when the physical changes to observe during specimen
28 testing do not involve particle breakage but rather contact attrition.

29

30 Discrete-element method (DEM) based simulations are used to complement and extend difficult
31 experimental work (e.g. Liu *et al.*, 2023; Shi *et al.*, 2022; Wu *et al.*, 2022; Phan *et al.*, 2021). DEM
32 model results could then facilitate continuum model calibration. Also, if properly formulated, DEM
33 models may be scaled up to directly represent problems of engineering interest (e.g. Arroyo *et al.*
34 *et al.*, 2011; Ciantia *et al.*, 2016a; Zhang *et al.*, 2019; Cerfontaine *et al.*, 2021), bypassing continuum
35 modelling altogether. DEM may be used like a continuum phenomenological model, calibrating
36 microscale parameters on specimen scale responses. However, DEM models are more attractive
37 when information acquired at particle scale is leveraged to reproduce specimen-scale
38 phenomena.

39

40 A number of studies have explored time effects in granular soils using DEM. Kuhn & Mitchell
41 (1992, 1993), presented a 2D visco-frictional contact model, in which the tangential force acting
42 at contacts relaxed in time with a rate that was dependent on mobilized contact friction at the
43 same contact. The model was justified using rate process theory (RPT), by reference to atomic-
44 scale interactions at solid contacts (silica-silica bonds), activated through thermal energy. The
45 model achieved a strain rate decay similar to that of soils, while also recovering the influence of
46 mobilized macroscopic strength on the onset of creep failure. The RPT model was later
47 generalized to 3D (Kwok & Bolton, 2010), combined with a time-independent particle breakage
48 model (Liu *et al.*, 2019) or even applied to simulate rock creep (Gutierrez *et al.*, 2020).

49

50 Other researchers (Wang *et al.*, 2008; Wang *et al.*, 2014; Tong & Wang, 2015) have used a variety
51 of rheological models including viscous dashpots to describe contact interaction in DEM models
52 of sand. The behaviour predicted by these models is similar to that achieved with RPT, although
53 the contact model parameters are openly recognized as phenomenological and can only be
54 determined by examining their effect on specimen scale response.

55

56 The physical basis claimed for RPT DEM models is frequently lost during model calibration. This
57 is due to the difficulties inherent in using DEM to simulate long-duration experiments (creep
58 observations in the laboratory typically last hours or days; aging in field conditions is observed for
59 months and years). Dynamic DEM computation is advanced explicitly in time. Computational

60 stability requirements limit the time step to values that, in most circumstances, are well below 1
61 μs (Otsubo *et al.*, 2017b). This makes the computational load of any realistic DEM creep
62 simulation overwhelming, if the simulated time is fully reproduced in the simulation time.

63

64 Models based on RPT have bypassed this difficulty by means of material scaling, in which some
65 of the micromechanical model parameters are recalibrated to accelerate creep. For models based
66 on RPT this implies scaling up the viscous parameter value (Kuhn & Mitchell, 1992; Kwok &
67 Bolton, 2010) or, equivalently, to directly scale time (Gutierrez *et al.*, 2020). The scaling factor
68 applied is calibrated to match experimentally observed strain rate levels (Liu *et al.*, 2019; Gutierrez
69 *et al.*, 2020). The resulting scaling factors are very large numbers (for instance 10^{10} in Kuhn &
70 Mitchell) that dwarf the effect on the model viscous parameter of the physical variables
71 (temperature, activation energy,...) that would otherwise determine its value. It is then very difficult
72 to verify if the first-principles viscous parameter value selected is actually relevant for the material
73 at hand.

74

75 An interesting alternative to RPT is offered by models trying to explain soil creep through sub-
76 critical crack growth (Atkinson, 1984), as fracture propagation has an inherent time scale. This
77 was noted by Oldecop & Alonso (2007) in connection with rockfill time-dependent deformation.
78 DEM models of soil creep based on sub-critical crack growth have two potential advantages. One
79 is that fracture mechanics parameters relevant to geomaterials might be measured, allowing
80 microscale calibration. The second is that if grain breakage is included, the model output may be
81 also verified against laboratory measurements of GSD evolution. However, these two potential
82 advantages of fracture-based creep DEM models have not been yet fully exploited.

83

84 Kwok & Bolton (2013) showed how creep results similar to those of soils could be obtained from
85 a DEM model based on fracture. They were using bonded agglomerates to model grains, a
86 technique that severely curtails the number of grains that may be represented in the model and
87 no attempt was made to compare grain size distribution outcomes with laboratory results.
88 Fracture affected the bonds between agglomerate sub-particles and the parameters controlling
89 bond strength degradation were calibrated on specimen-scale responses. The same model was

90 later applied by Xu *et al.*, (2018) who showed its ability to reproduce non-isotach behaviour
91 (Tatsuoka *et al.*, 2008; Lade *et al.*, 2007). Calibration of bond strength degradation was done at
92 specimen scale. Some DEM GSD curves during creep were obtained, but not compared to any
93 laboratory result.

94

95 Charles (1958) power law was implemented into DEM to simulate the mechanical behaviour of
96 rockfill by Tapias *et al.*, (2015) and Alonso & Tapias (2019). Grains were represented using
97 breakable agglomerates, although fracture evolved at the agglomerate scale. The model was
98 complex to calibrate, using a mixture of microscale and specimen scale information, like yield
99 stress. Predictions of triaxial compression, including grain size distribution, were contrasted with
100 experiments, with varying success. As detailed by Tapias (2016), the dynamic computational
101 timestep was directly mapped into real time, assuming that every timestep during a creep phase
102 represented one second in the laboratory. Since the DEM computational timestep is not
103 independent of contact stiffness or particle size, such an approach is difficult to generalize,
104 particularly when the simulation involves breakage or stress-induced stiffening.

105

106 Xu *et al.*, (2018) used a different computational strategy: alternate or on-off computation. In this
107 approach most of the simulated process is advanced with a simplified ageing model and the full
108 dynamic model – run without any time or material scaling- is only switched-on when required to
109 recover equilibrium. This on-off computational strategy had been previously applied in some DEM
110 (Tran *et al.*, 2009) or FEM-DEM models (Ma *et al.*, 2015) of rockfill fracture-induced time evolution.
111 It has also strong analogies with the high-cycle continuum models (Niemunis *et al.*, 2005) applied
112 to evaluate the effect of long-term cyclic loads.

113

114 Building on that previous work this study presents a new DEM model based on fracture to simulate
115 creep in sands. The model conjugates conceptual simplicity, computational efficiency and ease
116 of calibration. For the first time a DEM model for sand creep is proposed with the ability to match
117 laboratory experiments in terms of creep strain, creep strain rates, GSD evolution and particle
118 breakage. Thanks to the on-off computational strategy applied, the computational load is

119 moderate even for very long real-time experiments, while still using physically based values of
 120 DEM material parameters.

121

122 **2. Model description**

123 **2.1 Rough contact model**

124 In this work we use only spherical discrete elements. A simplified Hertz-Mindlin frictional contact
 125 model (e.g. Thornton, 2015), is modified to take into account roughness effects on the normal
 126 stiffness component. The modification follows Otsubo *et al.*, (2017a), who proposed a model
 127 (Figure 1) with three successive regimes to describe the influence of particle roughness on
 128 contact normal stiffness.

129

130 In this model when $\delta \leq \delta_{T1}$, the contact response is in the asperity-dominated regime and:

$$131 \quad F_n = F_{nT1} \left(\frac{\delta}{\delta_{T1}} \right)^c \quad (1)$$

132 when $\delta_{T1} < \delta < \delta_{T2}$, the contact response is in the transitional-dominated regime:

$$133 \quad F_n = F_{nT2} \left(\frac{\delta - \delta_1}{\delta_{T2} - \delta_1} \right)^b \quad (2)$$

134 when $\delta_{T2} \leq \delta$, the contact response is in the Hertzian regime:

$$135 \quad F_n = \frac{4}{3} E' \sqrt{r'} (\delta - \delta_1 - \delta_2)^{1.5} \quad (3)$$

136 where F_n is the normal contact force. E' and r' are calculated by Equation (4), (5):

$$137 \quad E' = \left(\frac{1 - \nu_1^2}{E_1} + \frac{1 - \nu_2^2}{E_2} \right)^{-1} \quad (4)$$

$$138 \quad r' = \left(\frac{1}{r_1} + \frac{1}{r_2} \right)^{-1} \quad (5)$$

139 In the previous equations δ is the particle overlap, E_1 and E_2 are young's modulus of the two
 140 contacting particles; ν_1 and ν_2 their Poisson's ratios; r_1 and r_2 their radii.

141

142 δ_{T1} , δ_{T2} are critical contact overlaps for three transition regimes that correspond to transitional
 143 contact normal forces of F_{nT1} and F_{nT2} . These are functions of roughness S_q and elastic
 144 parameters. δ_1 and δ_2 are both also functions of roughness S_q :

$$145 \quad \delta_1 = n_1 S_q \quad (6)$$

$$146 \quad \delta_2 = n_2 S_q \quad (7)$$

147 where n_1 and n_2 are model parameters. Parameters b and c are constants that ensure stiffness
148 continuity, and both depend only on $\delta_1, \delta_2, \delta_{T1}, \delta_{T2}$. It is apparent that when $S_q = 0$, a standard
149 Hertzian contact model is recovered. More details about the model may be found in Otsubo *et al.*,
150 (2017a).

151

152 **2.2 Particle failure model**

153 *2.2.1 Time-independent failure*

154 The time-independent particle failure model employed was introduced by Ciantia *et al.*, (2015)
155 and has been used for several studies at the specimen scale (Ciantia *et al.*, 2016b; 2019a; 2019b).
156 The model was later refined by Zhang *et al.*, (2021) to take into account the rough contact model
157 just introduced.

158

159 The model is inspired by analytical studies of Russell & Muir Wood (2009), who combined a two-
160 parameter material strength criterion with the analysis of the elastic stress distribution induced by
161 point loads on a sphere to obtain a failure criterion for the loaded particle. The limit condition was
162 expressed as:

$$163 \quad \kappa_{mob} \leq \kappa \quad (8)$$

164 where κ_{mob} and κ are the mobilised and intrinsic strengths of particles. Mobilized strength, in turn,
165 is given by:

$$166 \quad \kappa_{mob} = f(\chi, \nu) \frac{F_n}{\pi R^2 \sin^2 \theta} \quad (9)$$

167 where F_n is the normal force acting at a contact, R particle radius and θ the solid contact angle
168 'seen' from the centre of the sphere (Figure 2). The function $f(\chi, \nu)$ –given in Russell & Muir Wood
169 (2009)– expresses a condition of maximum tensile stress in the sphere derived from the elastic
170 stress distribution and strength criteria. The strength criteria (Christensen, 2004) is formulated
171 using parameters χ and κ , but these are directly related to σ_c and σ_t the uniaxial compressive and
172 tensile strengths of the material. Russell & Muir Wood (2009) worked out analytically that this
173 failure criterion remains valid under multiple contact loads; Tapias *et al.*, (2015) obtained the same
174 result using numerical simulation.

175

176 For application in DEM models, Ciantia *et al.*, (2015) observed that equations (8) and (9) imply a
 177 condition on contact forces and that such condition could be expressed as the product of a
 178 material property, the characteristic particle limiting stress, σ_{lim} , and the force contact area, A_F

$$179 \quad F_n \leq \frac{\kappa}{f(x,v)} \pi R^2 \sin^2 \theta = \sigma_{lim} A_F \quad (10)$$

180 This later realization led to generalize the Russell & Muir Wood (2009) failure condition. From the
 181 point of view of the material, the generalization introduced (Ciantia *et al.*, 2015) was double. First,
 182 a particle size-effect was introduced in the failure criteria, and second, randomness was
 183 introduced in the strength assignment for particles of a single size, to indirectly represent the
 184 influence of aspects such as particle shape on particle breakage. This generalized particle limiting
 185 stress is implemented through an assignment of particle strength given by

$$186 \quad \sigma_{lim} = \sigma_{lim,0} \left(\frac{d}{d_0} \right)^{-\frac{1}{m_p}} (1 + X_{0,1} var) \quad (11)$$

187 where d is the element diameter, $\sigma_{lim,0}$ is the mean value of strength at the reference diameter
 188 d_0 , and m_p a material parameter. The value var is the coefficient of variation of the distribution of
 189 particle strength for particles of diameter d_0 , assumed normal. The symbol $X_{0,1}$ represents a
 190 random number sampled from the standard normal distribution. All the parameters relevant to the
 191 crushing model may be obtained from single particle crushing tests.

192

193 From the point of view of the force contact area, the relevant expression for the rough Hertzian
 194 contact model is

$$195 \quad A_F = \pi r' \delta \quad (12)$$

196 With the meaning of r' and δ as given in the previous section.

197

198 Once failure takes place, a particle will split into 14 balls using a splitting scheme described in
 199 Ciantia *et al.*, (2015). The sibling particles are oriented so that the local z-axis is aligned with the
 200 normal component of the maximum contact force (see inset in Figure 3). After that, the assignment
 201 of particle strength σ_{lim} for 14 balls will also be carried out. To ensure computational efficiency, a
 202 numerical comminution limit is imposed to stop crushing for particles smaller than a certain
 203 diameter, d_c .

204 The grading state index, I_G , (Muir Wood, 2007; Figure 3) is used to quantify grading evolution. It
 205 is given (Figure 3) by the area ratio of current grain size distribution (GSD) curve to a fractal limit
 206 GSD with fractal factor $\beta = 2.6$:

$$207 \quad \frac{M_{(L<d)}}{M_T} = \frac{d^{3-\beta} - d_{min}^{3-\beta}}{d_{max}^{3-\beta} - d_{min}^{3-\beta}} \quad (13)$$

208 where $M_{(L<d)}$ is the mass of particles whose diameter smaller than d ; M_T is the total mass. d_{max}
 209 and d_{min} are maximum and minimum diameter for the sample (Table 1).

210

211 A fraction of the broken particle volume is lost upon breakage; it is assumed that the material lost
 212 corresponds to fines and those are accounted for in post-processing, to refine estimates of
 213 material grading evolution. The fractal distribution (Equation 13) is also used to estimate the GSD
 214 of the mass lost at each particle splitting event, using as d_{max} the smallest particle generated
 215 during the event. Previous studies, (Ciantia *et al.*, 2015; 2016a; 2019a; 2019b; Zhang *et al.*, 2021),
 216 have shown that the amount of volume lost at the specimen scale using the 14 particle split is
 217 small, and that increasing the number of siblings does not significantly modify macroscale model
 218 results.

219

220 2.2.2 Time-dependent failure

221 To introduce a time dependency in the failure model Charles law (1958) is used. As noted by
 222 Alonso & Tapias (2019), Charles law is simply an empirical description of experimental
 223 observations of crack growth under tensile loading, and takes the form,

$$224 \quad v = v_0 \left(\frac{K}{K_c} \right)^n \quad (14)$$

225 where v is crack velocity, v_0 is a reference velocity, K_c represents material toughness, K the
 226 stress intensity factor and n is the stress corrosion index.

227

228 Following Broek (1986), the stress intensity factor K for mode I (tensile) failure can be expressed
 229 as

$$230 \quad K = \beta_c \sigma_I \sqrt{\pi a} \quad (15)$$

231 where β_c is a parameter dependent on the geometry of crack and cracked body, σ_I is the far field
 232 applied tensile stress and a is the crack half-length. Toughness K_c corresponds to the stress

233 intensity at critical conditions, arrived at by increasing far field stress, crack length or both. If σ_t is
 234 the tensile strength, leading to uncontrolled crack growth for the initial fracture geometry, it results
 235 that

$$236 \quad \frac{K}{K_c} = f_k \frac{\sigma_t}{\sigma_t} \quad (16)$$

237 where f_k is a geometry dependent term that will be characteristic of the test employed to measure
 238 toughness, of specimen size, temperature, etc. Russell & Wood (2009) note that in the
 239 Christensen material model tensile strength σ_t and limit strength κ and are proportional, so that

$$240 \quad \sigma_t = \frac{\sqrt{3}}{(1+\chi)^2} \kappa = \frac{\sqrt{3}}{(1+\chi)^2} f(\chi, \nu) \sigma_{lim} = f'(\chi, \nu) \sigma_{lim} \quad (17)$$

241 They also show that the maximum elastic tensile stress along the diameter beneath a contact
 242 force is proportional to applied contact stress, so that

$$243 \quad \sigma_t = f_d(\nu) \frac{F_n}{A_F} = f_d(\nu) \sigma_{mob} \quad (18)$$

244 where σ_{mob} is the applied contact stress. It turns out that, for a point loaded sphere, with a crack
 245 aligned with the load and located at the maximum of elastic tensile stress,

$$246 \quad \frac{K}{K_c} = f_k \frac{f_d(\nu) \sigma_{mob}}{f'(\chi, \nu) \sigma_{lim}} \quad (19)$$

247 This result is merely indicative for real sand grains, given the limitations of the material model,
 248 and the variability in grain shape and in the location and nature of flaws within the grain.
 249 Nevertheless, it does suggest that a reasonable time-dependent behaviour may be obtained if
 250 Charles law is applied in the DEM model simply through:

$$251 \quad v = v_0 \left(\frac{\sigma_{mob}}{\sigma_{lim}} \right)^n \quad (20)$$

252 where σ_{mob} the maximum normal contact stress acting on a particle, and σ_{lim} is the particle
 253 strength; v_0 and n are the same as those in Eq. (14). Charles law is a model for crack growth, so
 254 the elements need to be seeded with initial cracks, of half-length a_0 . The crack half-length a grows
 255 in time as:

$$256 \quad a = a_0 + v \Delta t \quad (21)$$

257 where v is crack velocity, and Δt is time interval for updating the crack. To model crack
 258 propagation in the DEM model, the approach proposed by Tapias *et al.* (2016) is used. A virtual
 259 crack half-length a is treated as a particle internal variable evolving with time following eq. (21).
 260 Upon sample generation a value of a_0 uniformly distributed in the range $0.001d$ to $0.5d$ is

261 assigned to each particle in the model. To limit the complexity of the model this approach does
262 not aim to represent any realistic crack geometry, disregarding crack orientation and implicitly
263 assuming a virtual crack to start from the centre of the sphere and develop radially in two
264 diametrically opposite directions. Once the virtual crack length is equal to the particle diameter (a
265 is equal to particle radius) the particle will fail and is replaced by the 14 particle arrangement as
266 done for the time independent criterion. Upon breakage initial crack half-lengths a_0 are also
267 randomly assigned to every sibling particle.

268

269 **2.3 Computational strategy**

270 In this work an off-DEM ageing technique is employed to advance the simulation during creep
271 test phases. As shown in Figure 4 in this technique the dynamic DEM computation stages
272 alternate with periods of off-DEM ageing. During a dynamic computation stage the discrete model
273 runs with all its features activated: elements and boundary walls move, contacts are created or
274 lost, contact forces and contact stresses change, assigned particle cracks grow and particles are
275 broken if and when a breakage criterion is attained.

276

277 During off-DEM ageing periods only crack growth is active. Crack growth velocity during this
278 phase is assumed constant, given by the mobilized contact stress determined at the end of the
279 previous dynamic computation step. As time advances during the off-ageing period the crack
280 growth mechanism would lead to failure in some particles; such particles are identified and
281 counted, until their number attains a certain pre-specified limit value, (n_{age}).

282

283 The dynamic DEM computation is resumed, breaking all the n_{age} particles. This creates a dynamic
284 disruption that is computed until equilibrium is again attained until the specified creep stress state.
285 After some sensitivity analyses (Lei *et al.*, 2023) the value of n_{age} was established as 30 (0.26%
286 of total particle number) to limit the initial disruption. The dynamic computation step is finished
287 when stress fluctuation around the target is stabilized and a minimum number of cycles (150.000)
288 has elapsed. All creep simulation phases start and end with a dynamic computation stage.

289

290 **3. Model calibration for Fontainebleau sand**

291 **3.1 Previously calibrated Fontainebleau sand parameters**

292 Fontainebleau sand is a quartz sand widely used in geotechnical research, which has been used
293 as model material in previous DEM studies (Ciantia *et al.*, 2015; Ciantia *et al.*, 2019a; Zhang *et*
294 *al.*, 2021). The parameters calibrated by Zhang *et al.*, (2021) for Fontainebleau sand using the
295 rough-crushable model are reported in Table 1. It is worth noting that incorporating surface
296 roughness enables the model to capture the initial softer response observed in experimental data
297 (e.g. Wong & Coop, 2023) whilst using a material shear modulus G very similar to that of real
298 quartz sand particles. This is different from previous DEM studies of quartz sand creep (Kwok &
299 Bolton, 2013; Liu *et al.*, 2019), where much lower G values were adopted to capture realistic
300 single particle force displacement curves.

301

302 Particle rotation was inhibited to roughly simulate the interlocking effects due to non-spherical
303 particle shapes. This is a computationally efficient simplification (Ting *et al.*, 1989; Calvetti, 2008;
304 Arroyo *et al.*, 2011) that may be seen as a limit case for classical rolling-resistance contact models
305 (Rorato *et al.*, 2021).

306

307 **3.2 Crack growth parameters**

308 The two crack growth parameters to calibrate are the reference velocity v_0 and the stress
309 corrosion index n . Following, Tapias *et al.* (2015; 2019) and in line with fracture growth data from
310 by Oldecop & Alonso (2007), in this work, $v_0 = 0.1$ m/s was used. Whilst for rockfill material the
311 presence of water is known to influence the stress corrosion index (Oldecop & Alonso, 2007), the
312 effect of water on creep on quartz sands has been shown to be negligible (Leung *et al.*, 1997;
313 Olson *et al.*, 2002). To calibrate n , literature data for glass, synthetic and natural quartz, and
314 quartz rich sandstone is presented in Figure 5 and $n = 60$ was hence selected as a pragmatic
315 choice for this study of quartz sands. It's worth mentioning that for rocks in general, the stress
316 corrosion index is highly sensitive to both ambient factors (such as stress level, temperature,
317 presence and chemistry of the pore fluid) (Brantut *et al.*, 2013) and compositional factors. Both
318 limestones and clay-rich sandstones are highly sensitive to water presence (Olson *et al.*, 2002;
319 Nara *et al.*, 2012).

320

321 **4. Model application**

322 **4.1 DEM model for element tests**

323 The rough contact model and the time-independent particle failure model were implemented using
324 the C++ plug-in capability of PFC 3D (version 5.00.40; Itasca, 2017). Time dependent failure and
325 the off-DEM simulation advance algorithm were implemented using FISH, the high-level
326 programming language of PFC.

327

328 Following Ciantia *et al.*, (2019a), a representative cubic volumetric element (REV) of 4 mm side,
329 containing 11500 particles, was formed using the radius expansion method. The particle size
330 distribution (0.1 to 0.4 mm) was selected to match closely that of Fontainebleau NE34 sand. The
331 target initial void ratio was set to 0.65 which would correspond to a relative density of 65% for the
332 sand. Boundaries of the REV cube were set as rigid walls; wall motion was servo-controlled.

333

334 The model implementation was verified using results from a series of oedometer loading tests on
335 Fontainebleau NE34 sand presented by Ciantia *et al.*, (2019a). Figure 6 shows the results. The
336 newly implemented DEM model agrees well with the oedometer loading curve and with the
337 grading evolution deduced from the laboratory test series.

338

339 **4.2 Creep under oedometric conditions**

340 Brzesowsky *et al.* (2014) report results of a 2-day creep experiment on vacuum dried quartz sand
341 (with diameter $d=378\pm 22\mu m$) under oedometric conditions at 21.7 MPa. A creep test under the
342 same conditions was simulated using the parameters calibrated for Fontainebleau sand. During
343 the creep phase the constant vertical stress was enforced by continually adjusting top wall
344 position using a stress-controlled scheme.

345

346 The simulation results are shown in Figure 7. Numbered red dots mark the episodes of dynamic
347 computation; the time in between is off-DEM ageing time. Figure 7 shows that the simulated creep
348 curve appears very similar to the experimental creep trend, even if no effort was made to adjust
349 parameters to match the test results. Note that most model parameters are not specific of
350 Fontainebleau sand. As detailed in previous work (Ciantia *et al.* 2019a; Zhang *et al.* 2021) most

351 parameter values employed in the model were selected considering generic properties of quartz-
352 dominated sands and other geomaterials (e.g. glass beads, sandstone). That applies to the elastic
353 parameters (G , ν), to the parameters controlling immediate breakage (m_p , $\sigma_{lim,0}$, var) and to the
354 parameters describing roughness effects on contact stiffness (S_q , n_1 , n_2). Also, and perhaps most
355 importantly, the parameters controlling fracture growth have been evaluated with reference to
356 other quartz materials (see section 3.2). The only aspects specifically matched to Fontainebleau
357 sand properties are the grain size distribution and the value of interparticle friction μ , selected to
358 match a triaxial test result (see Ciantia *et al.* 2019a).

359

360 The properties of the Fontainebleau sand and the quartz sands from the Heksenberg Formation
361 tested by (Brzesowsky *et al.*, 2014) are presented in Table 3. As the model scales particle strength
362 with particle size, the smaller grain size of Fontainebleau would partly explain the faster
363 accumulation of breakage-related deformation for Heksenberg sand visible in Figure 7. It is also
364 relevant that the void ratio before creep was slightly lower for the DEM model (at 0.584) than for
365 the lab test (at 0.615), as looser specimens are prone to undergo more breakage (e.g. Sohn &
366 Buscarnera, 2018).

367

368 Figure 8 shows the juxtaposed vertical stress record during the 15 dynamic computation stages
369 that took place during the oedometer creep phase. The simultaneous breakage enforced to start
370 every dynamic computation stage creates a small shock that is rapidly damped by the boundary
371 stress-control algorithm. It can be seen that the amplitude of the vertical stress fluctuation remains
372 below 4% of the target creep stress. Although the number of particles to break at the beginning
373 of every dynamic computation remains constant, the response recorded at the different dynamic
374 computation stages presents some variation.

375

376 **4.3 Creep under triaxial conditions**

377 To simulate creep under triaxial conditions the REV was first loaded isotropically up to a confining
378 stress of 10 MPa. From the oedometer test results (Figure 7) this stress level was expected to
379 guarantee significant particle breakage during shearing. A strain-controlled standard triaxial

380 compression path was then applied with top wall velocity slow enough (0.01 m/s) to avoid any
381 inertial effects, (resulting inertial number $I < 2.68 \times 10^{-4}$).

382

383 The test was initially taken to failure (at 30% deviatoric strain) to identify the available shear
384 strength. This corresponded to a maximum deviatoric stress q_{max} of 22.6 MPa. Ten creep tests
385 were then performed with deviatoric stress maintained constant at different fractions of the
386 maximum (q/q_{max} values between 0.2, and 0.9). In those tests the specimen was first loaded to
387 the desired level of mobilized strength and then deviatoric stress was maintained for at least 10^4
388 s, unless shear failure was attained before.

389

390 Some examples of the deviatoric stress evolution plots obtained in the triaxial tests are illustrated
391 in Figure 9. As in the oedometer, there are some dynamic oscillations around the deviatoric creep
392 stress, but they remain limited, smaller, for instance, than the oscillations due to continuous shear
393 at failure during the monotonic reference path.

394

395 Figure 10 shows the volumetric strain versus axial strain plots for the creep phase. Shearing creep
396 strains are accompanied always by volumetric contraction. This contractive creep was also
397 observed by Karimpour & Lade (2013) when testing Virginia Beach sand at high confining
398 pressures (8 MPa). Creep strains are more contractive at lower deviatoric stress levels than at
399 higher ones, as observed by Lv *et al.* (2017) in their laboratory tests of quartz and coral sands.

400

401 Figure 11 shows axial strain ε_a evolution during the simulated creep phases. Creep strain is linear
402 on log time when mobilized strength q/q_{max} lies below a certain limit (0.7 in this study). When
403 q/q_{max} exceeds this value, obvious inflection points can be observed, and creep strain
404 accumulates at faster rates leading to creep failure. This is classically described as tertiary creep
405 (Augustensen *et al.*, 2004) and has been observed in similar triaxial experiments on quartz sands
406 both at high confining stress (Karimpour & Lade, 2013; $q/q_{max} > 0.8$) and at low confining stress
407 (Murayama *et al.*, 1985; $q/q_{max} > 0.95$).

408

409 Before the onset of creep failure the axial strain rate $\dot{\epsilon}$ decreases in time. As shown in Figure
410 12(a) an approximate linear evolution can be observed when plotting the creep strain data in a
411 double log axis diagram. This is characteristic of primary creep. When creep takes place at higher
412 mobilized strength the strain rate remains briefly constant (secondary creep) and then increases
413 (tertiary creep). That sequence, however, is not always monotonous and some oscillations
414 between increasing and decreasing strain rates are visible in tests creeping close to shear failure
415 (q/q_{max} above 0.8). Similar oscillations were also present in laboratory triaxial creep tests at high
416 mobilised shear strength in friable Antelope Valley sand (Lade & Liu, 1998).

417

418 DEM simulation results at $q/q_{max}=0.4$ and $q/q_{max}=0.75$ were selected to compare with triaxial
419 creep results of Virginia beach sand at the same mobilized ratios (Karimpour & Lade, 2010; 2013).
420 The comparisons (Figure 12b) show good agreement with the experimental results during primary
421 creep. In the case of $q/q_{max}=0.75$, the strain rate tends to stabilize in both cases at the end of
422 creep although the tertiary creep phase appears earlier in the Fontainebleau DEM experiments.

423

424 It is customary to describe the strain rate evolution during creep using the m parameter defined
425 by Singh and Mitchell (1968) as:

$$426 \quad m = -\frac{\Delta \log \dot{\epsilon}}{\Delta \log t} \quad (22)$$

427 Linear regression was applied to obtain m values from the test data. Some example fits are
428 illustrated in (Figure 13). Following Augustesen *et al.* (2004), the fit is only applied to the primary
429 creep section. To compare with the existing experimental results, the strain rate (%/min) at 10
430 min ($\dot{\epsilon}_{10min}$) was also obtained. All the strain rate parameters are presented in Table 2.

431

432 The strain rate m parameter obtained from the DEM simulations is presented in Figure 14(a)
433 alongside previous laboratory and DEM simulation results. The m value obtained here is
434 practically independent of deviatoric creep stress (Figure 14a) as observed on the laboratory tests
435 of Toyoura quartz sand (Murayama *et al.*, 1984) and tailings sand (Mejia *et al.*, 1988).

436

437 This independence of m on stress level is also presented on the RPT-based 2D DEM simulations
438 of Kuhn & Mitchell (1993), who excluded creep rupture data from their fit. A similar trend is visible
439 –with less clarity- on the RPT-based 3D DEM simulation of Kwok & Bolton (2010). By way of
440 contrast the bond strength degradation model of Kwok & Bolton (2013) shows a strong reduction
441 of m as the mobilized strength at creep was increased. This comparison might be affected by the
442 criteria -not always explicit- employed by different authors to fit the m parameter.

443

444 A simpler comparison may be established in terms of creep strain rate magnitude. Figure 14(b)
445 shows its value evaluated 10 min after the initiation of creep. The results from this work fit nicely
446 within the range of the experimental data, with the more friable tailing materials showing faster
447 rates, the low-stress tests on quartz Toyoura showing slower creep rates and the high stress tests
448 on quartz Virginia beach sand closer to the simulations. Interestingly, by this measure the RPT
449 based simulations are those further off the experiments, whereas the bond-strength based Kwok
450 & Bolton (2013) model lies closer to the results presented in this study.

451

452 **4.4 Grading evolution during triaxial creep**

453 GSD curves during triaxial creep were obtained during the simulations. Figure 15 illustrates the
454 results for $q/q_{max}=0.2, 0.5, 0.7,$ and 0.9 . As mobilized strength at creep increases the shift of the
455 PSD curves is more significant.

456

457 The breakage-driven nature of the creep strains is also made transparent when the evolution of
458 the grading index with time is represented, as done in Figure 16. The similitudes with Figure 11
459 are notable, for instance in the inflection points noted at values of q/q_{max} above or equal to 0.75.

460

461 The amount of breakage during creep is expressed through the change in breakage index in
462 Figure 17. The results from the simulation are very similar to those measured after laboratory
463 experiments on Virginia beach sand by Karimpour & Lade (2013), although breakage
464 accumulation is somewhat faster in Virginia sand. As it was the case for the previous comparison
465 with Heksenberg sand, this discrepancy may be partly explained by the smaller grain size of
466 Fontainebleau (Table 3).

467

468 Note that, for the purpose of this comparison, four creep simulations were extended to last one
469 day, to match the creep period in the original lab experiments. Most of that period was simulated
470 as off-DEM ageing, whereas the time simulated using dynamic DEM computation (less than 0.1s)
471 was only a minimal fraction of that period. The computer running time (using a workstation with
472 an CPU of 12th Gen Intel(R) Core(TM) i9-12900K (3.20 GHz) and 32GB RAM) extended to 41h
473 for the most demanding case (that with q/q_{\max} of 0.85). The vast majority (>99,9%) of running time
474 was spent on the dynamic DEM computations. The practical advantages of the computational
475 strategy adopted are thus evident.

476

477 **5. Discussion**

478 The DEM model presented is highly idealized, in that some fundamental sand grain features such
479 as shape are only indirectly accounted for through particle strength variability, rotation blocking,
480 etc. Also, the link between roughness and contact friction has not been considered, nor the
481 possibility of inducing a low stress time-dependent mechanism through evolving roughness.
482 Despite those limitations it is likely that the model micromechanics might be usefully exploited to
483 gain insight into macroscopic features. As an example, Figure 18 presents the evolution of the
484 broken fraction of particles of certain size ranges, d_i . This fraction is obtained as the cumulative
485 number of particles of size d_i broken to time t divided by the current number of particles of size d_i
486 at time t . The graph indicates not just that larger particles break more –something that follows
487 from particle strength size dependency- but that tertiary creep is characterized by an increased
488 participation of larger particles on breakage, as the broken fraction increases faster with particle
489 size.

490

491 Laboratory compression experiments on relatively uniform and large glass beads (Takei *et al.*,
492 2001) showed jumps on creep strain associated with breakage events. In continuously graded
493 specimens DEM studies with unbreakable particles (e.g. Liu *et al.*, 2023) have shown that strong
494 force chains are preferentially channelled through larger particles. The non-monotonous increase
495 in strain rates observed during tertiary creep in Figure 11 may be thus related to the faster
496 breakage rate of the larger particles.

497

498 Another peculiarity of the model presented is that time-dependent breakage parameters were
499 calibrated using fracture data from larger quartz-dominated specimens. This was done for
500 pragmatic reasons, to avoid the experimental and statistical complexity of trying to measure
501 fracture growth on isolated sand grains. It is thus interesting to consider the results in Figure 19,
502 showing the effect of normalized creep stress on the time elapsed until the onset of the tertiary
503 creep phase (known as time to failure) in the simulations and in sandstone creep experiments.
504 Note that in the sandstone data creep and failure stress levels are net of the post-fracture residual
505 frictional stress (Brantut *et al.*, 2013). That is the stress level at which all cohesion is lost and
506 shear strength is purely frictional, which, by analogy, is taken as zero for sands. The comparison
507 is favourable and supports the relevance of sandstone subcritical crack growth measurements for
508 the understanding of sand creep.

509

510 **6. Conclusion**

511 This work has described a discrete element model to explain time effects in sands based on
512 subcritical crack growth. The model has been applied to study creep at large stress in quartz sand
513 and the results obtained compare favourably with available experimental evidence in terms of
514 creep strain, creep strain rates and grading evolution. Previous DEM models of sand creep had
515 not resulted in such a wide agreement with independent laboratory experimental work. The
516 continuous IG evolution during creep can be captured using this model, something which is almost
517 impossible to do in laboratory tests. An increased participation of larger particles on breakage
518 during tertiary creep was first made transparent, and this is significant to understand sand creep
519 behaviour. The model has been implemented using a simple but efficient on-off computational
520 strategy to get rid of the overwhelming computational load associated with long creep tests in
521 DEM. Importantly, this computational strategy does not interfere with material parameter
522 determination.

523

524 The model may be applied as it is to study some of the fundamental controls on sand creep, such
525 as relative density (Colliat-Dangus *et al.*, 1983) or initial grading (Karimpour & Lade, 2013) as well
526 as other time-related phenomena like stress relaxation and, including a suitable time-scaling

527 procedure, to examine the effect of variable strain rate. It is also easy to envisage some relevant
528 model generalizations: for instance, as it is based on a contact law that features grain roughness,
529 it may be also simply adapted to represent low-stress time-dependent contact maturing. The use
530 of models such as this will likely facilitate the study of time dependent phenomena in granular
531 materials.

532

533

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538

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721 Tables

722 Table 1 Calibration parameters for time-to-fracture rough-crushable model

G/GPa	ν	μ	particle failure criterion				Contact roughness			Crack propagation			
			m_p	$\sigma_{lim,0}/\text{GPa}$	var	d_s/d_{50}	d_{max}/mm	d_{min}/mm	$S_q/\mu\text{m}$	n_1	n_2	ν_0	n
32	0.19	0.275	12	3.75	0.38	0.55	0.27	0.01	0.6	0.05	5	0.1	60

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724 Table 2. Triaxial creep strain rate parameters

q/q_{max}	0.2	0.3	0.4	0.5	0.6	0.7	0.75	0.8	0.85	0.9
m	0.7713	0.7356	0.7903	0.7484	0.7353	0.7246	0.7360	0.7238	0.723	0.7251
$\dot{\epsilon}_{10min}$ (%/min)	0.00043	0.001	0.0017	0.0038	0.0086	0.0263	0.0558	0.1247	0.30	0.855

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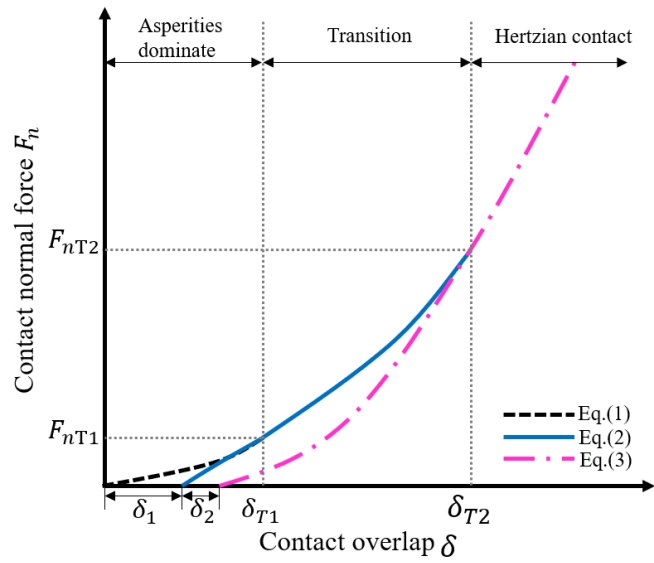
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Table 3. Basic properties of different quartz sands considered in this study

Property	Units	Virginia beach sand (Karimpour & Lade 2013)	Heksenberg Formation sand (Brzesowsky et al., 2014)	Fontainebleau sand NE 34 (Ciantia et al. 2019a)
Median grain size, d_{50}	mm	0.638	0.378	0.21
Uniformity coefficient		1.4	1.12	1.53
Min. void ratio, e_{min}		0.53	Unknown	0.51
Max. void ratio, e_{max}		0.759	Unknown	0.9
Specific gravity, ρ_s		2.65	2.65	2.65
Shape descriptor		subangular	subrounded	subangular

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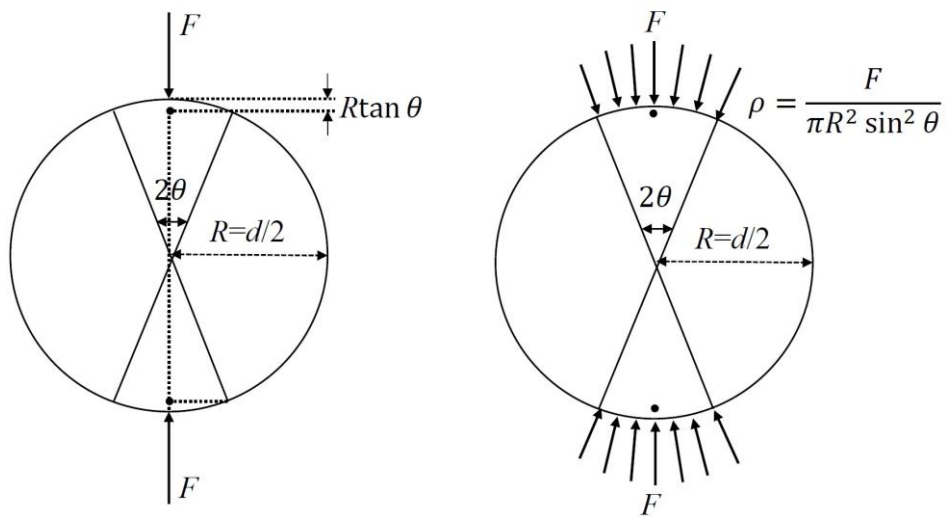
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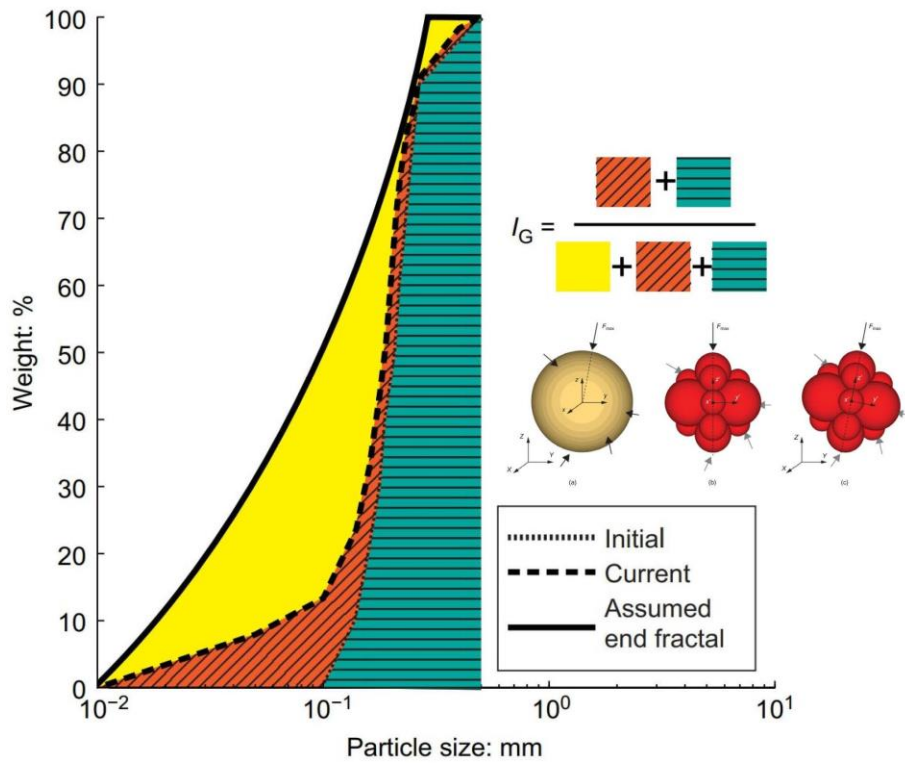
Figure 1. Rough surface normal contact model (adapted from Otsubo *et al.*, 2017a)



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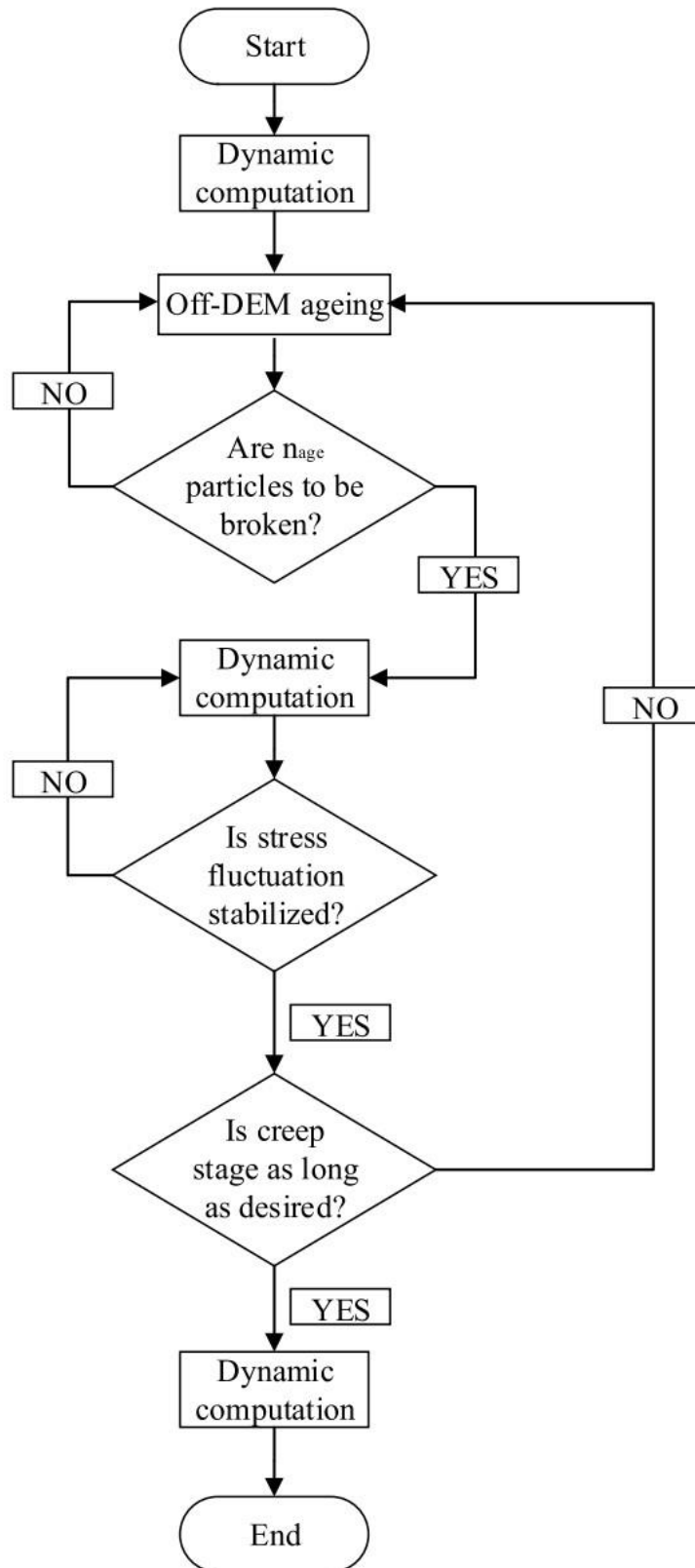
Figure 2. Contact geometry description in the failure model



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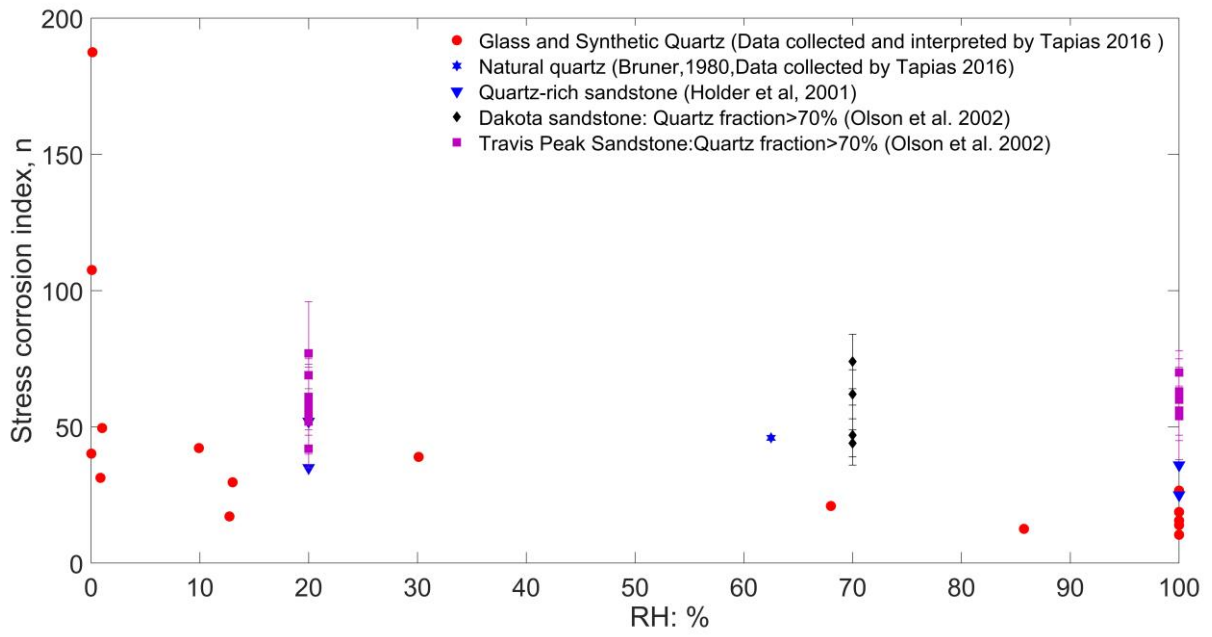
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Figure 3. Schematic definition of grading index I_G and (inset) particle split model



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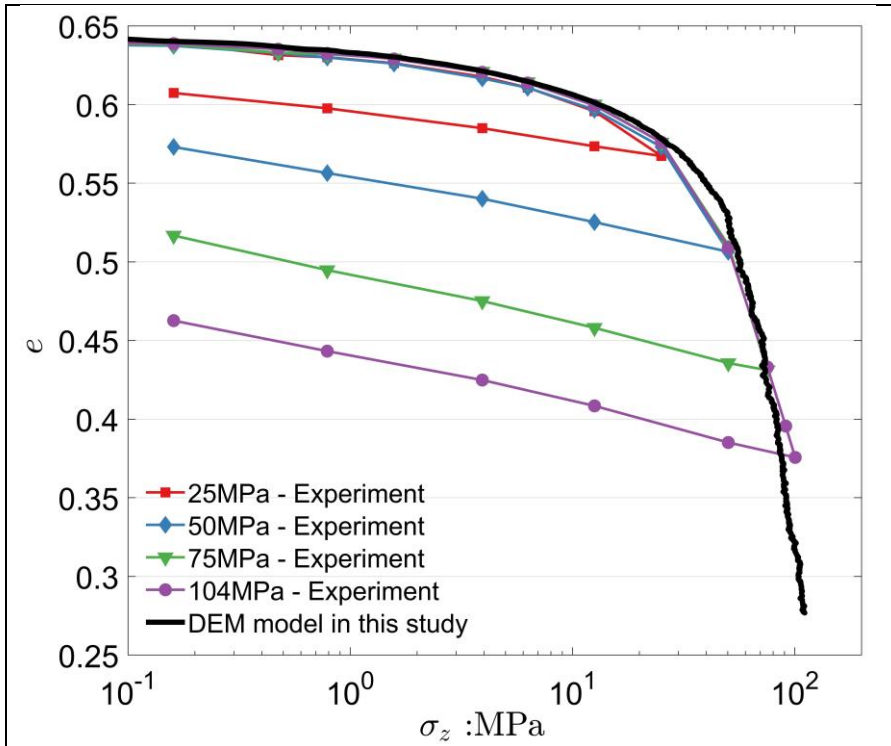
Figure 4. Flow chart for on-off DEM computation during creep phases



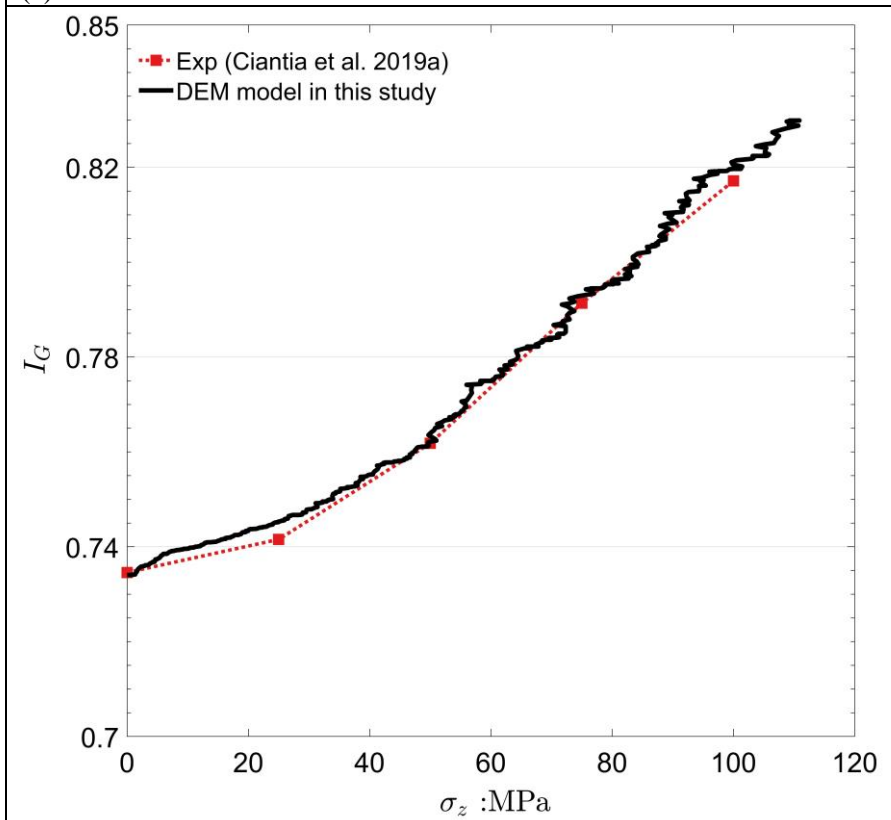
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Figure 5. Stress corrosion index data for some materials relevant to quartz sands

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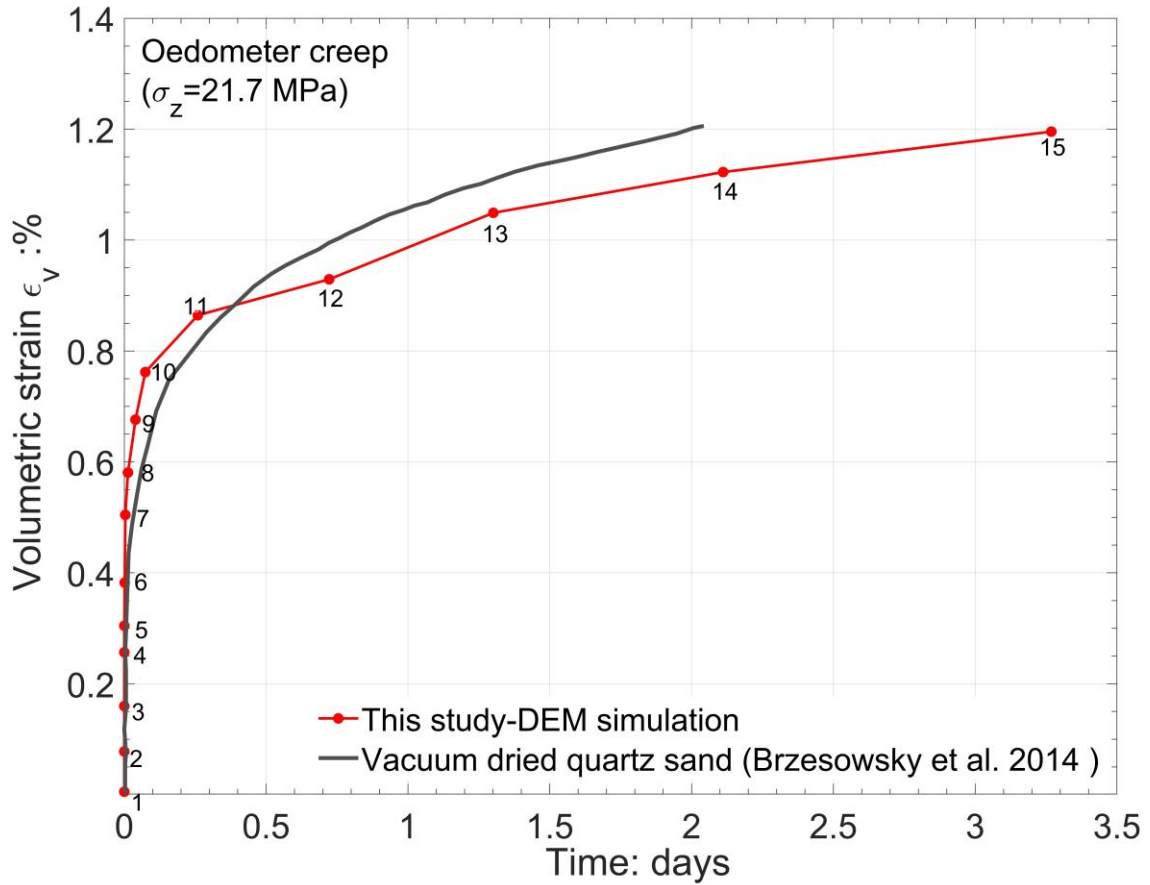
(a)



(b)

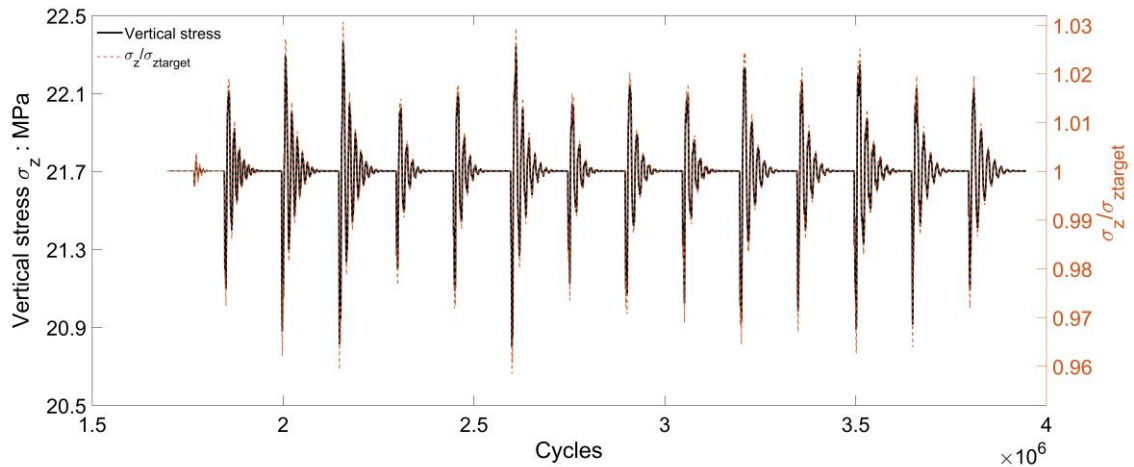
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Figure 6. Oedometer test calibration against experiments reported by Ciantia *et al.* (2019a): (a) void ratio e vs vertical stress σ_z (b) I_G evolution with σ_z



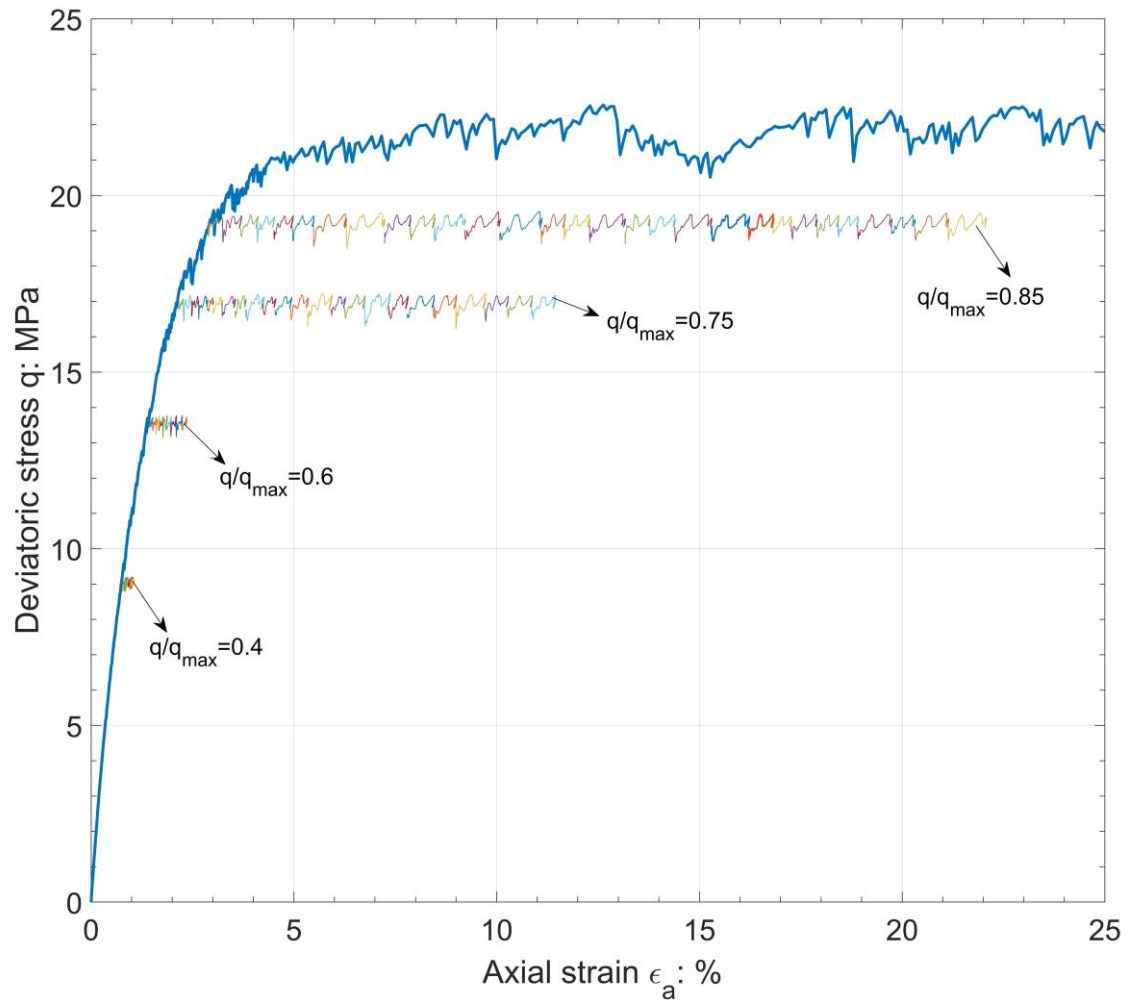
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Figure 7. Oedometric creep: simulation for the calibrated Fontainebleau sand and experimental result for vacuum dried quartz sand (Brzesowsky *et al.*, 2014). Dots in the simulation curve correspond to dynamic computation stages.



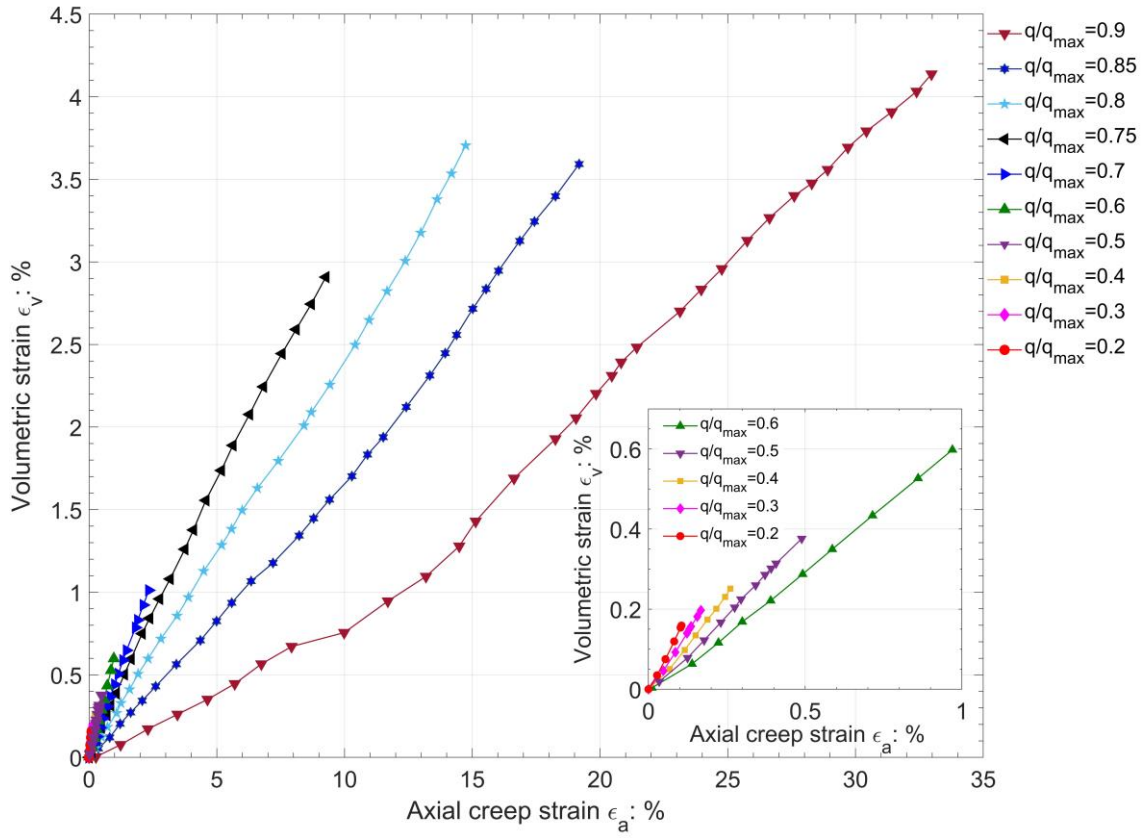
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Figure 8. Vertical stress fluctuations during the dynamic computation stages in the creep phase of the oedometric test. Left hand axis: absolute value. Right hand axis: normalized by target stress



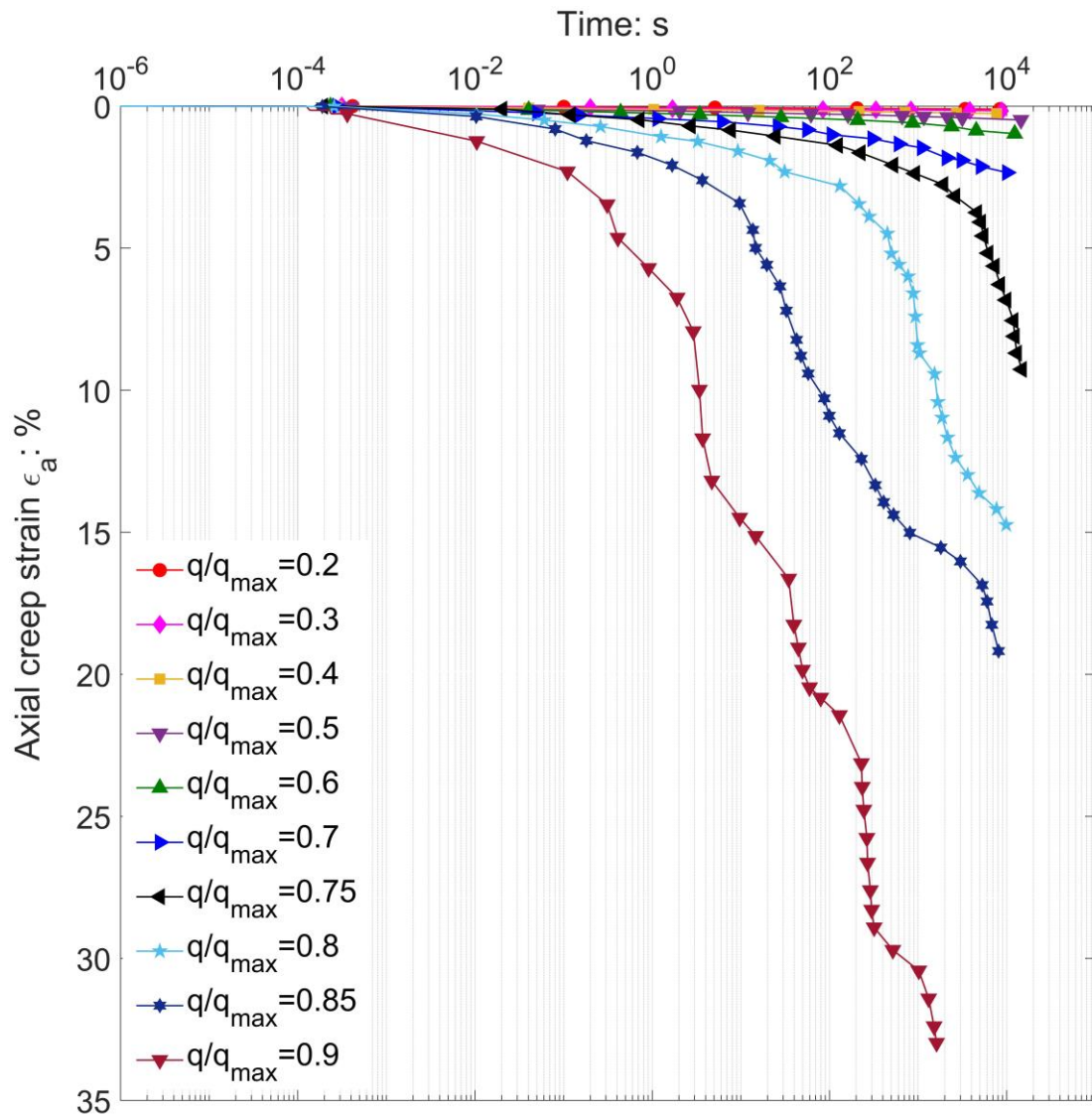
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Figure 9. Monotonic triaxial loading and example simulated triaxial creep curves



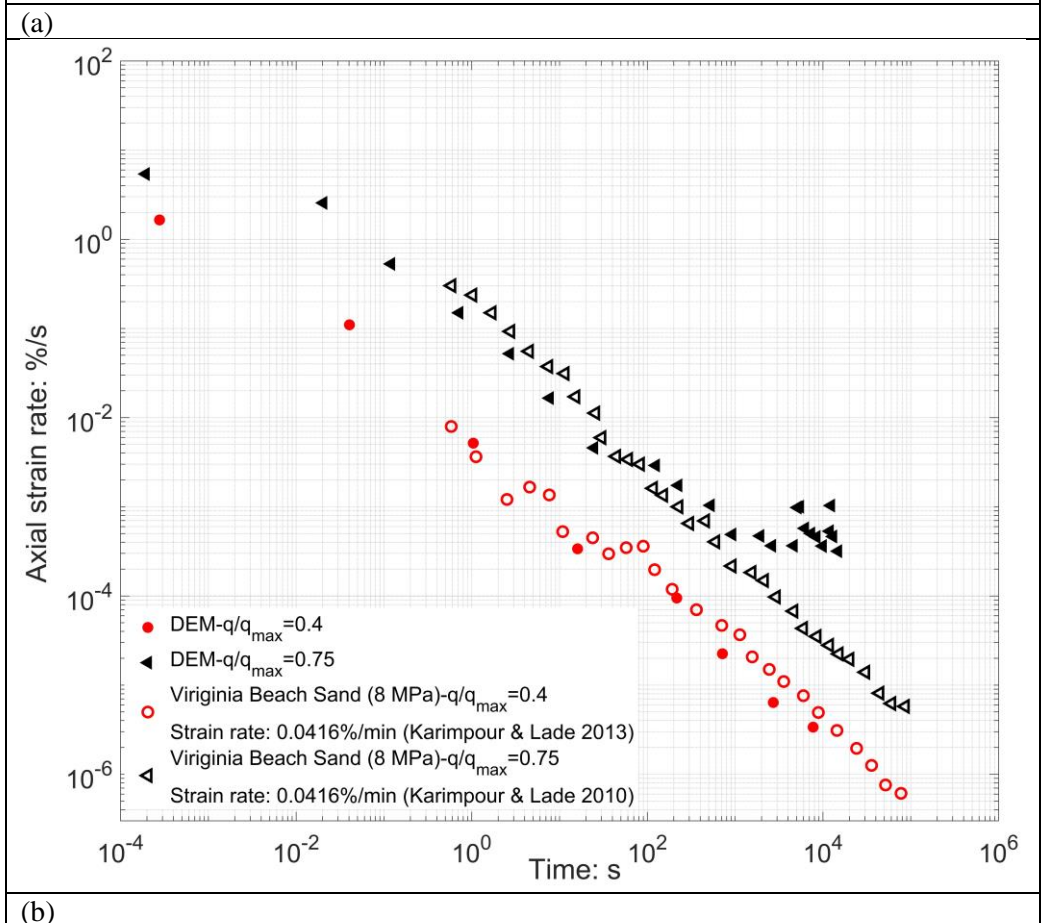
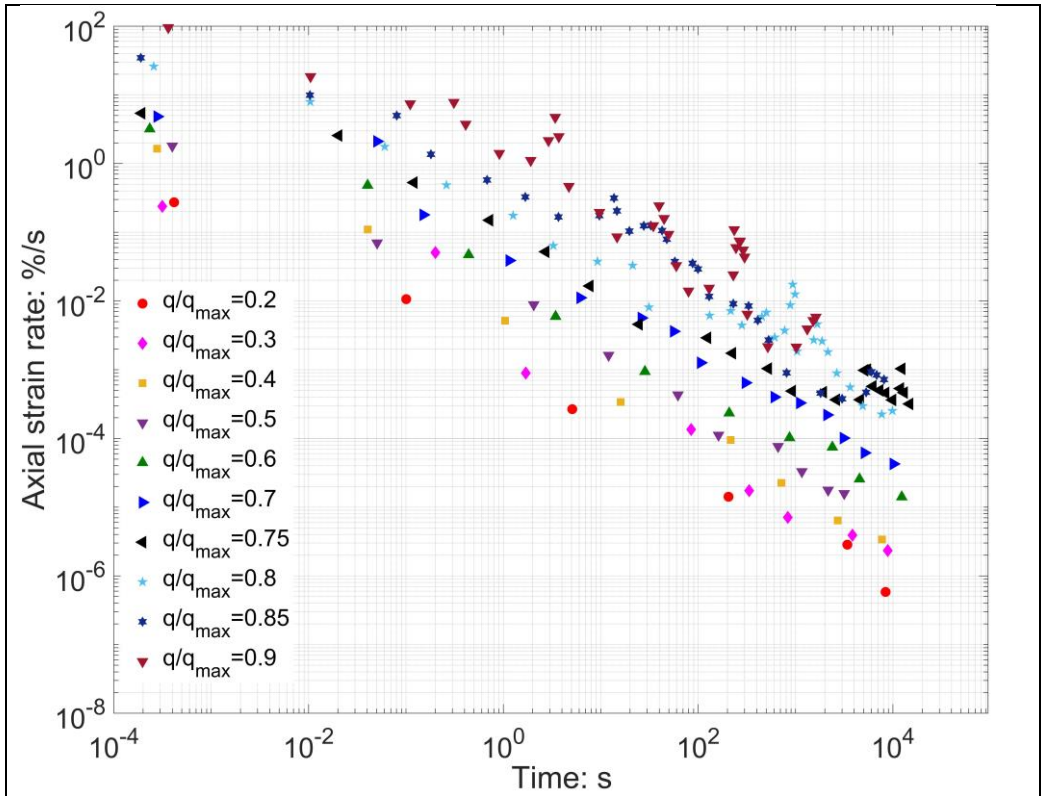
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Figure 10. Volumetric strain versus axial strain during triaxial creep



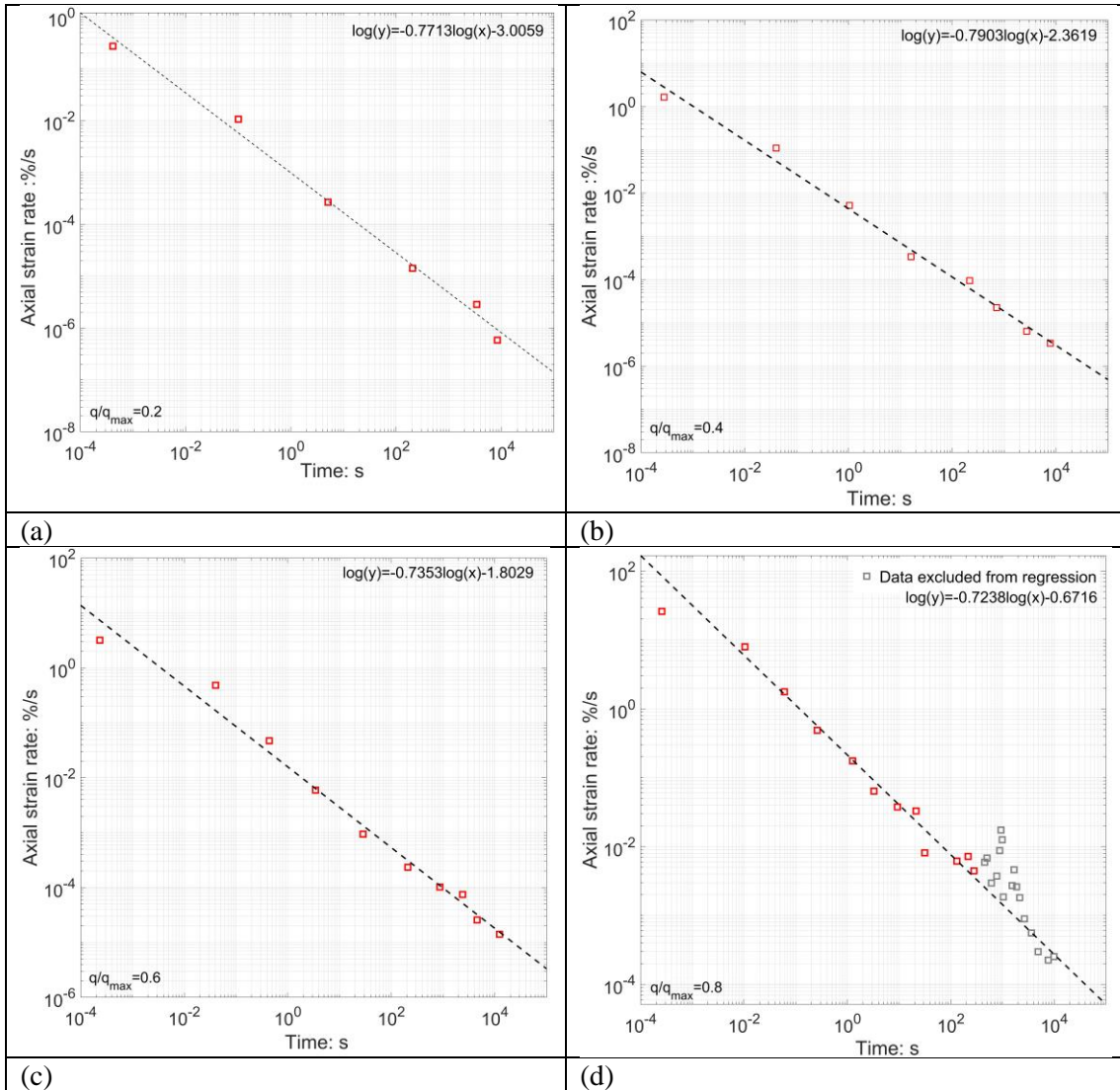
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Figure 11. Axial strain evolution during the creep phase of the simulations

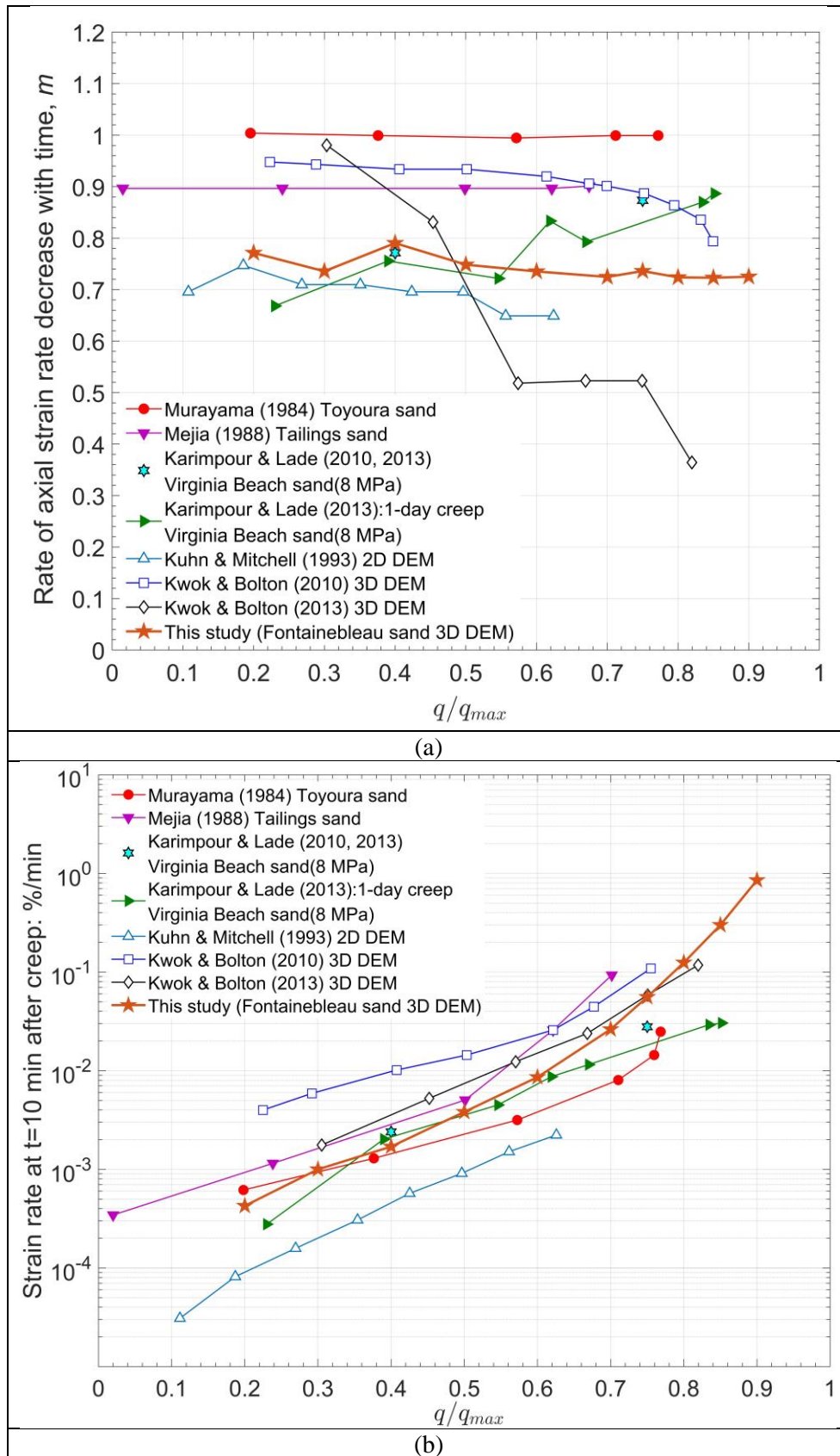


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773 Figure 12. Axial strain rate $\dot{\epsilon}_a$ versus time (a) DEM simulation results (b) Experimental comparison



774 Figure 13. Linear fitting of results from simulated creep triaxial stages to obtain the strain rate
 775 evolution parameter m defined by Singh and Mitchell (a) $q/q_{max} = 0.2$ (b) $q/q_{max} = 0.4$ (c) $q/q_{max} =$
 776 0.6 (d) $q/q_{max} = 0.8$

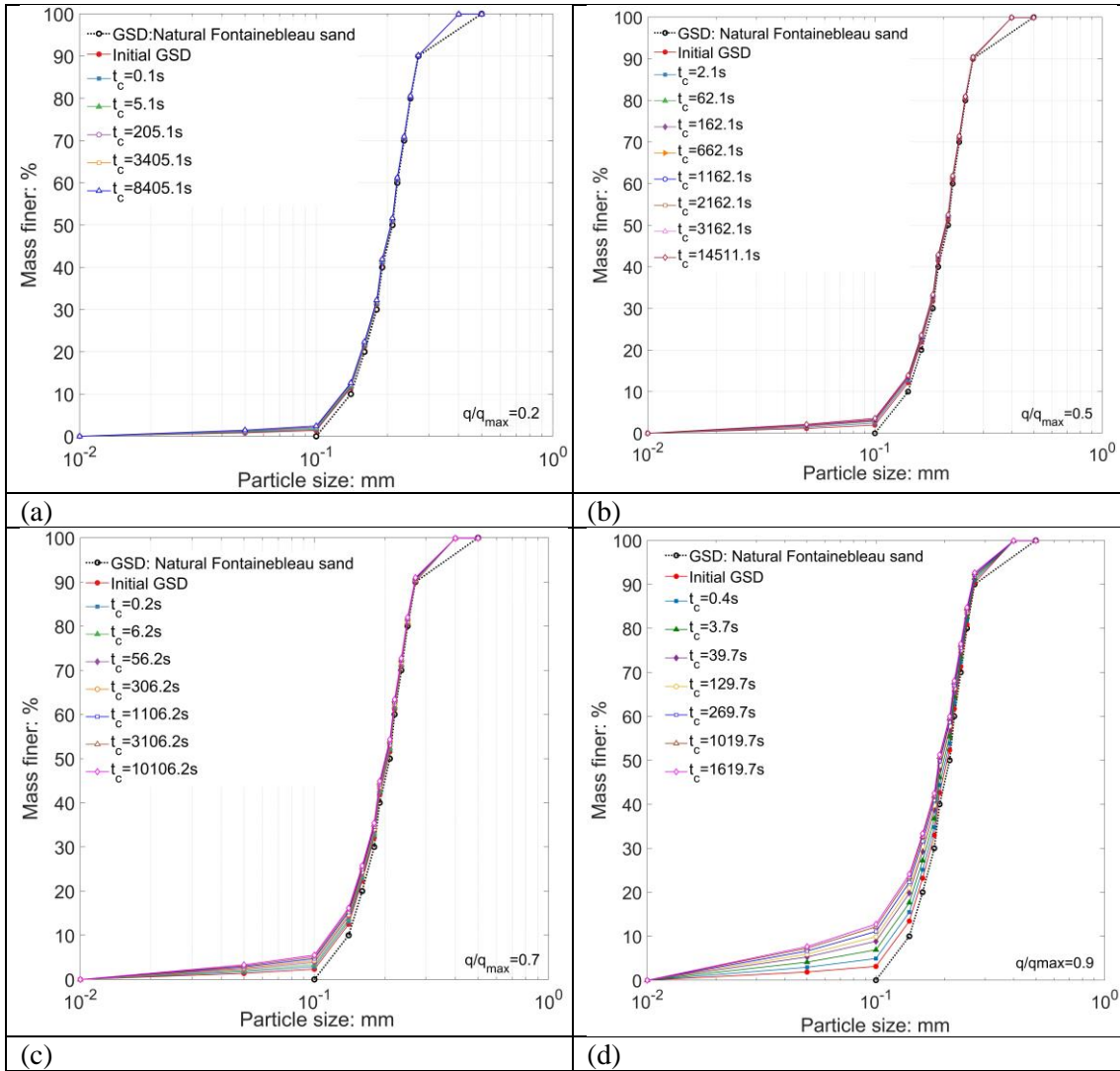


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Figure 14. Comparison of experimental and simulated results on axial strain rate (a) decrease with log time (b) value after 10 min creep



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Figure 15. Example of computed GSD evolution during triaxial creep (t_c : creep time)

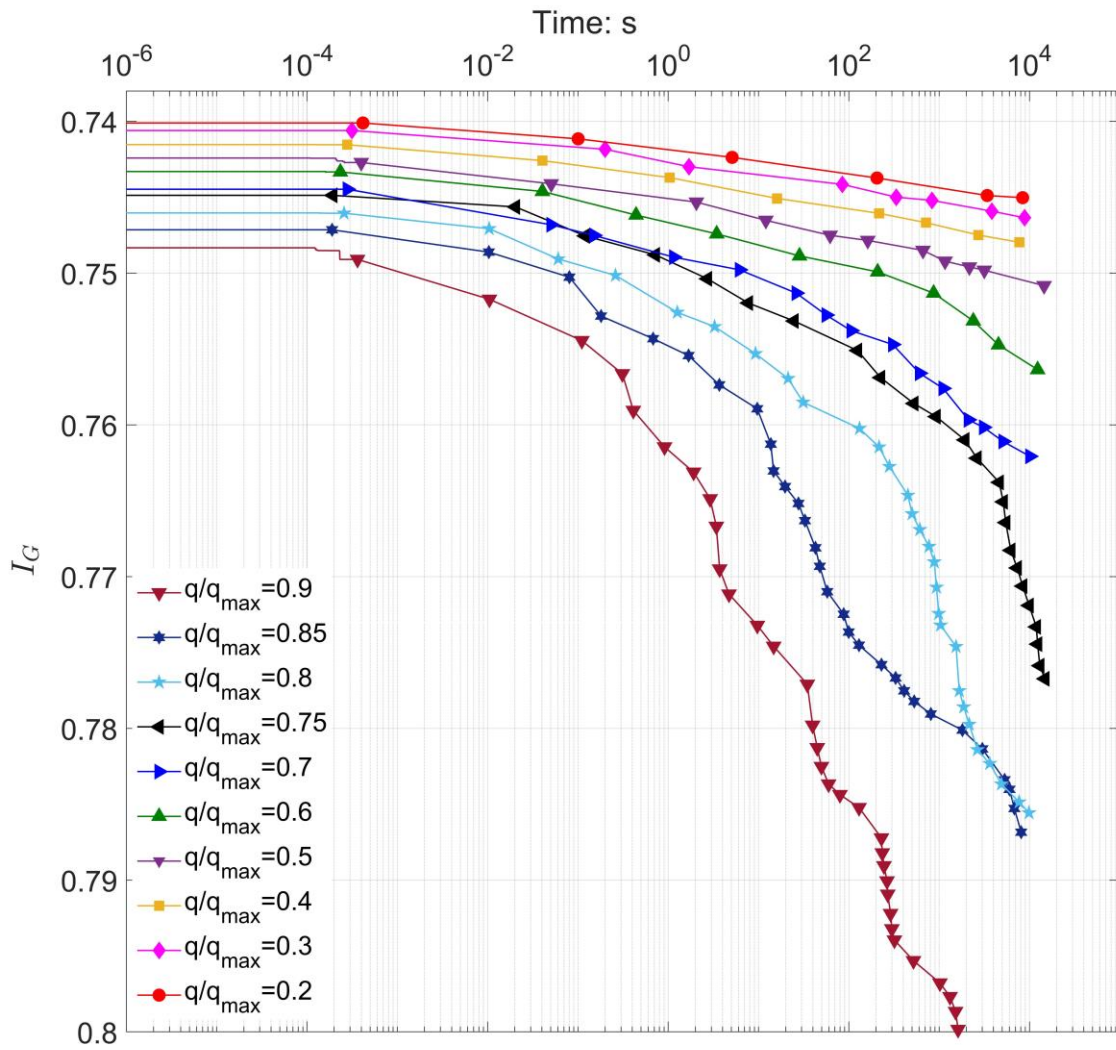
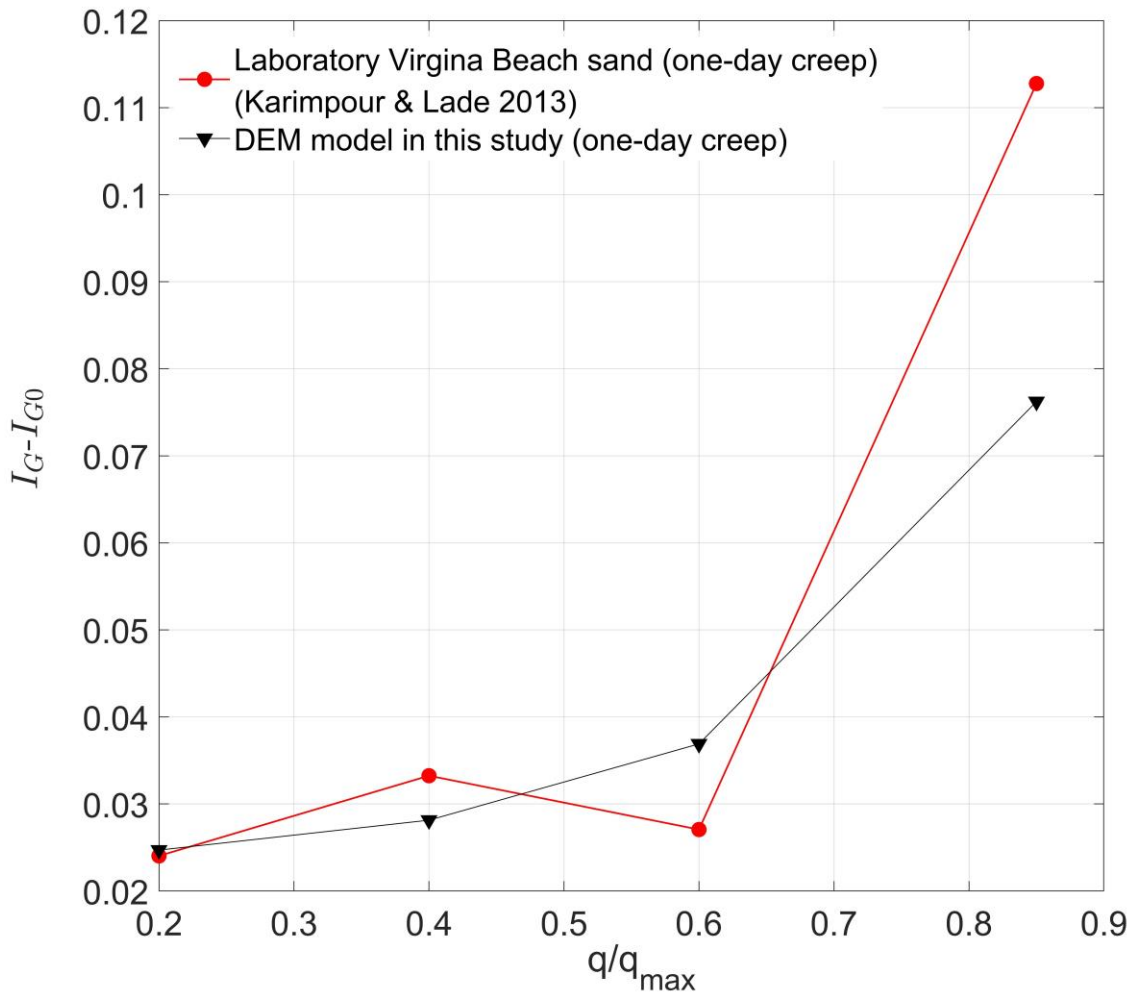


Figure 16. I_G evolution of Fontainebleau sand during triaxial creep.

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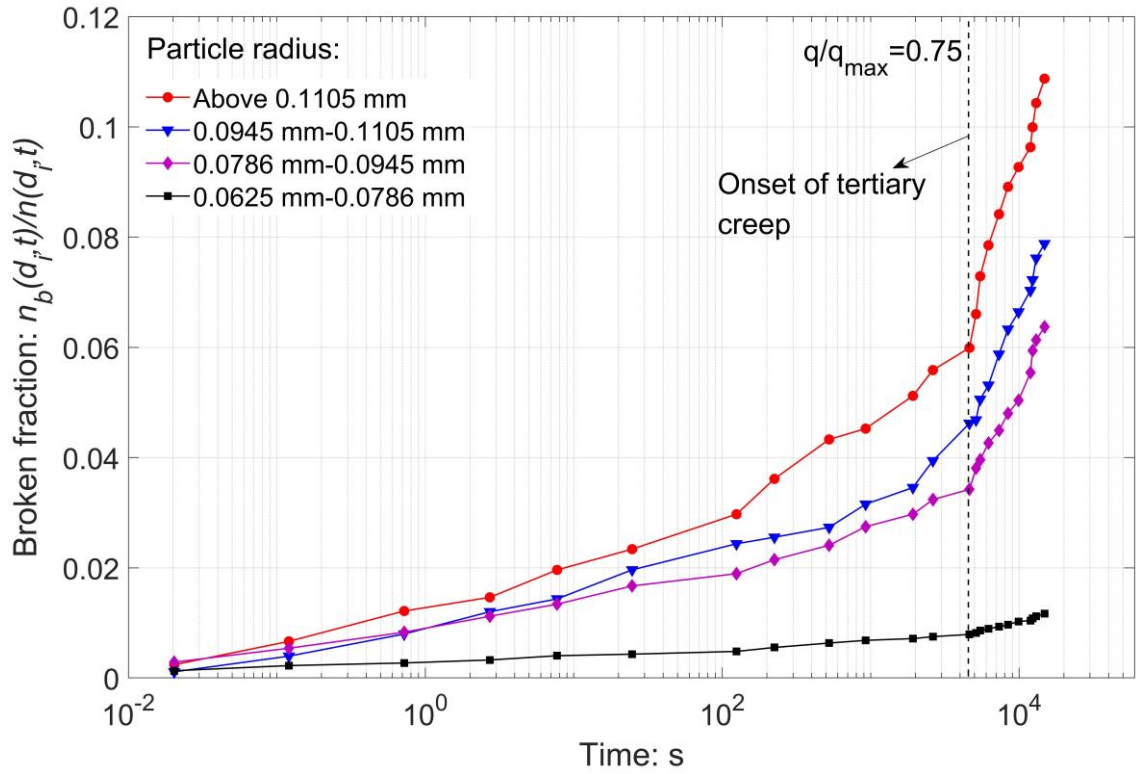


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784 Figure 17. Creep-induced change in breakage index I_G as a function of mobilized shear strength

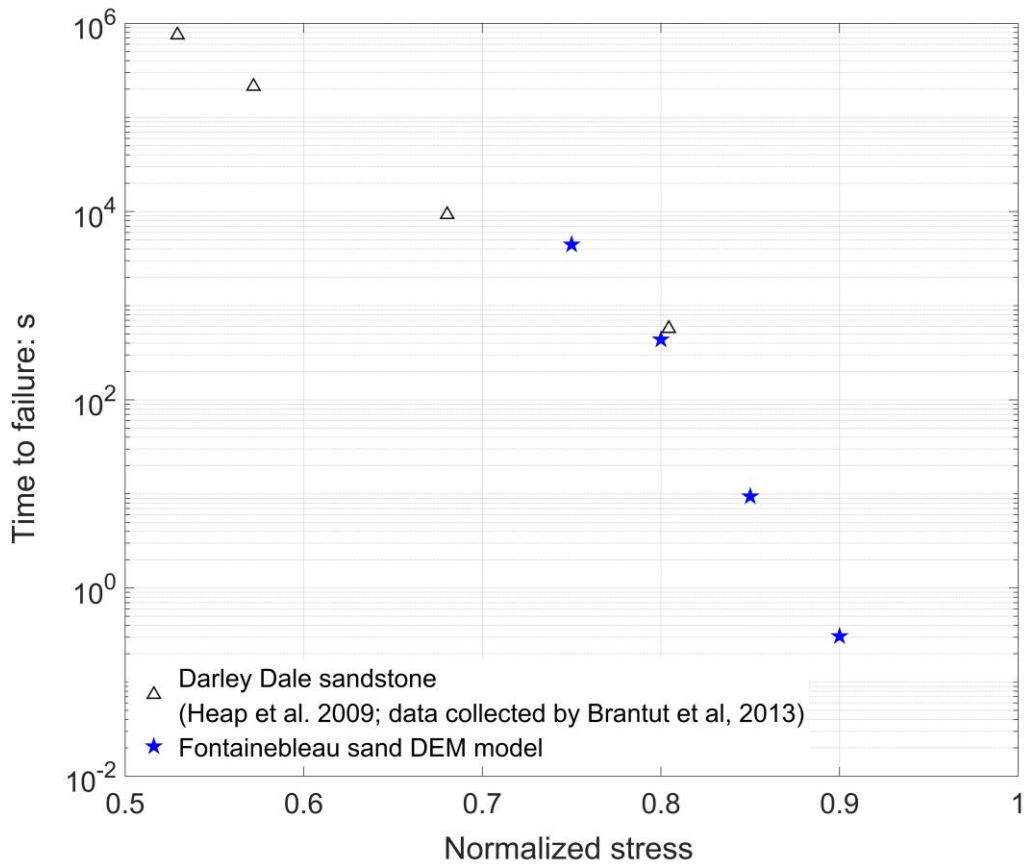
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during creep



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Figure 18. Evolution of the fraction of broken particles for different particle sizes



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Figure 19. Effect of normalized creep stress on time to failure (onset of tertiary creep) for quartz sandstone and simulated quartz sand